

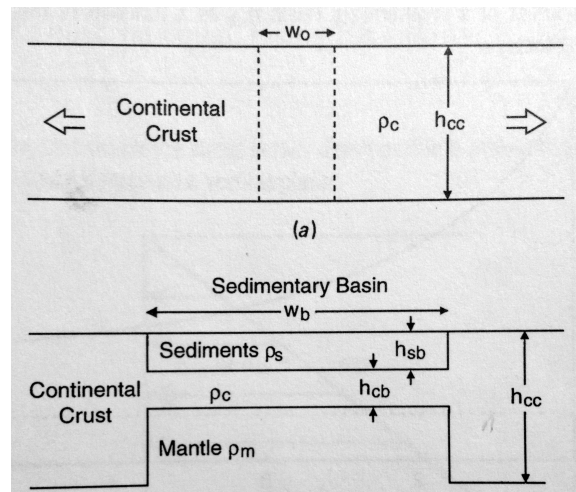
Earth Deformation Homework 1

Michał Dichter

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Problem 1 (T+S Problem 2-5)

We assume the setup of Figure 2-4 from Turcotte and Schubert:



We are given the following values:

$$h_{cc} = 35 \text{ km}$$

$$h_{sb} = 7 \text{ km}$$

$$\rho_m = 3300 \text{ kg/m}^3$$

$$\rho_{cc} = 2700 \text{ kg/m}^3$$

$$\rho_s = 2450 \text{ kg/m}^3$$

We need to solve equation (2-10):

$$h_{sb} = h_{cc} \left(\frac{\rho_m - \rho_{cc}}{\rho_m - \rho_s} \right) \left(1 - \frac{1}{\alpha} \right)$$

Plugging in the values above we get:

$$\alpha \equiv \frac{w_b}{w_0} \approx 1.4$$

Problem 2 (Rotational symmetry of a crystal)

Let us take some matrix A :

$$A = \begin{pmatrix} a & b & c \\ d & e & f \\ g & h & i \end{pmatrix}$$

Matrix of a 90-degree counter-clockwise rotation about x :

$$R_x(90) = \begin{pmatrix} 1 & 0 & 0 \\ 0 & 0 & -1 \\ 0 & 1 & 0 \end{pmatrix}$$

Matrix of a 90-degree counter-clockwise rotation about y :

$$R_y(90) = \begin{pmatrix} 0 & 0 & 1 \\ 0 & 1 & 0 \\ -1 & 0 & 0 \end{pmatrix}$$

Matrix of a 90-degree counter-clockwise rotation about z :

$$R_z(90) = \begin{pmatrix} 0 & -1 & 0 \\ 1 & 0 & 0 \\ 0 & 0 & 1 \end{pmatrix}$$

The matrix A rotated 90 degrees about x :

$$A_{\text{rot.x}} = R_x A R_x^{-1} = \begin{pmatrix} a & -c & b \\ -g & i & -h \\ d & -f & e \end{pmatrix}$$

The matrix A rotated 90 degrees about y :

$$A_{\text{rot.y}} = R_y A R_y^{-1} = \begin{pmatrix} i & h & -g \\ f & e & -d \\ -c & -b & a \end{pmatrix}$$

The matrix A rotated 90 degrees about z :

$$A_{\text{rot.z}} = R_z A R_z^{-1} = \begin{pmatrix} e & -d & -f \\ -b & a & c \\ -h & g & i \end{pmatrix}$$

The problem states that we must have:

$$A_{\text{rot.x}} = A_{\text{rot.y}} = A_{\text{rot.z}} = A$$

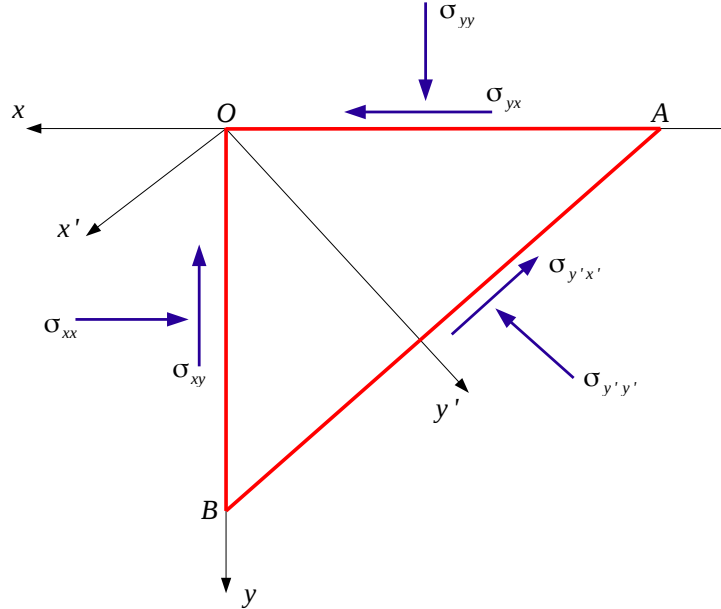
Thus we can conclude that all of the off-diagonal components must be zero and that $a = e = i$ such that the matrix A must be of the form:

$$A = \begin{pmatrix} a & 0 & 0 \\ 0 & a & 0 \\ 0 & 0 & a \end{pmatrix}$$

We therefore conclude that the matrix A has only *one* independent component!

Problem 3 (T+S Problem 2-13)

We can construct the following diagram:



Force balance in the \hat{x} direction tells us that:

$$\sigma_{yx}OA - \sigma_{xx}OB - \sigma_{y'x'} \cos(\theta)AB + \sigma_{y'y'} \sin(\theta)AB = 0$$

We can divide both sides by AB and rearrange:

$$\sigma_{y'x'} \cos \theta - \sigma_{y'y'} \sin \theta = \sigma_{yx} \cos \theta - \sigma_{xx} \sin \theta$$

Now multiply both sides by $\sin \theta$:

$$(\star) \quad \sigma_{y'x'} \sin \theta \cos \theta - \sigma_{y'y'} \sin^2 \theta = \sigma_{yx} \sin \theta \cos \theta - \sigma_{xx} \sin^2 \theta$$

Force balance in the \hat{y} direction tells us that:

$$\sigma_{yy}OA - \sigma_{xy}OB - \sigma_{y'x'} \sin(\theta)AB - \sigma_{y'y'} \cos(\theta)AB = 0$$

We can divide both sides by AB and rearrange:

$$\sigma_{y'x'} \sin \theta + \sigma_{y'y'} \cos \theta = \sigma_{yy} \cos \theta - \sigma_{xy} \sin \theta$$

Now multiply both sides by $\cos \theta$:

$$(\star\star) \quad \sigma_{y'x'} \sin \theta \cos \theta + \sigma_{y'y'} \cos^2 \theta = \sigma_{yy} \cos^2 \theta - \sigma_{xy} \sin \theta \cos \theta$$

Now use the symmetry of the stress tensor to say that $\sigma_{xy} = \sigma_{yx}$ and subtract (\star) from $(\star\star)$:

$$\sigma_{y'y'} (\sin^2 \theta + \cos^2 \theta) = \sigma_{xx} \sin^2 \theta + \sigma_{yy} \cos^2 \theta - 2\sigma_{xy} \sin \theta \cos \theta$$

Now recall that $\sin^2 \theta + \cos^2 \theta = 1$ and $2 \sin \theta \cos \theta = \sin 2\theta$. So we see that:

$$\boxed{\sigma_{y'y'} = \sigma_{xx} \sin^2 \theta + \sigma_{yy} \cos^2 \theta - \sigma_{xy} \sin 2\theta}$$

Problem 4 (Stress tensor)

We are given the following stress tensor:

$$\sigma = \begin{pmatrix} 1 & 1 & 0 \\ 1 & 1 & 0 \\ 0 & 0 & 2 \end{pmatrix}$$

We can calculate the pressure p :

$$p \equiv \frac{\sigma_{xx} + \sigma_{yy} + \sigma_{zz}}{3} = \frac{1 + 1 + 2}{3} = \frac{4}{3}$$

So we know the isotropic part of σ :

$$\sigma_{\text{iso}} = \begin{pmatrix} \frac{4}{3} & 0 & 0 \\ 0 & \frac{4}{3} & 0 \\ 0 & 0 & \frac{4}{3} \end{pmatrix}$$

And the deviatoric part of σ :

$$\sigma_{\text{dev}} = \sigma - \sigma_{\text{iso}} = \begin{pmatrix} -\frac{1}{3} & 1 & 0 \\ 1 & -\frac{1}{3} & 0 \\ 0 & 0 & \frac{2}{3} \end{pmatrix}$$

The eigenvalues of σ are:

$$\begin{aligned} \sigma_{\text{I}} &= 2 \\ \sigma_{\text{II}} &= 2 \\ \sigma_{\text{III}} &= 0 \end{aligned}$$

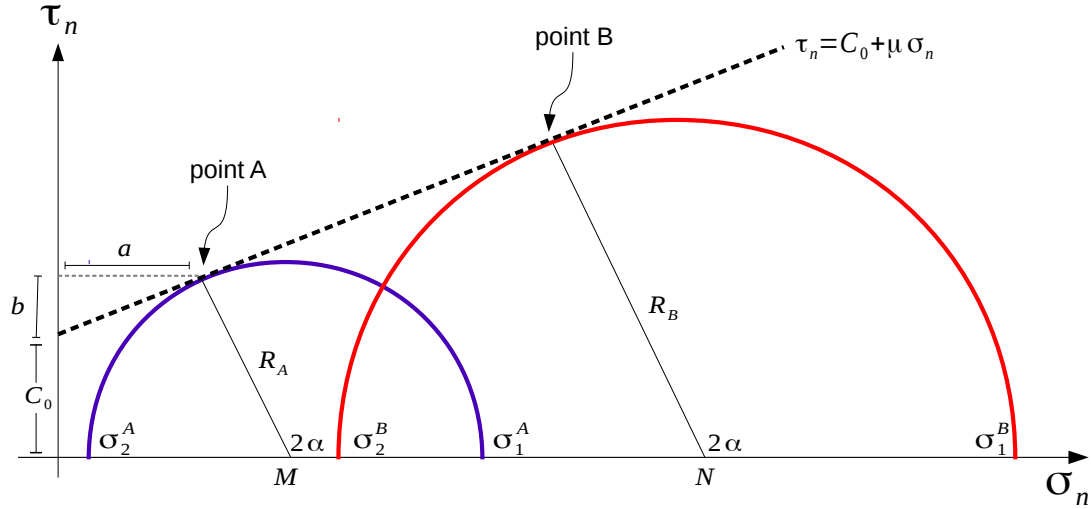
And the corresponding eigenvectors are:

$$\begin{aligned} \text{I} &= \begin{pmatrix} 0 \\ 0 \\ 1 \end{pmatrix} \\ \text{II} &= \frac{1}{\sqrt{2}} \begin{pmatrix} 1 \\ 1 \\ 0 \end{pmatrix} \\ \text{III} &= \frac{1}{\sqrt{2}} \begin{pmatrix} 1 \\ -1 \\ 0 \end{pmatrix} \end{aligned}$$

It is easy to see that these three eigenvectors are all mutually orthogonal (the dot product of any two equals zero) — they form an “orthonormal basis.”

Problem 5 (Mohr circle)

We can construct the following schematic diagram:



We can calculate the slope μ from point A to point B as a function of the angle 2α :

$$(*) \quad \mu = \left(\frac{\Delta\tau_n}{\Delta\sigma_n} \right)_{A \rightarrow B} = \frac{(R_B - R_A) \sin 2\alpha}{(\sigma_2^B - \sigma_2^A) + (R_B - R_A)(1 + \cos 2\alpha)}$$

We also know that the line perpendicular to $\tau_n = C_0 + \mu\sigma_n$ drawn from M to A (or from N to B) has a slope of $-1/\mu$:

$$(**) \quad -\frac{1}{\mu} = \left(\frac{\Delta\tau_n}{\Delta\sigma_n} \right)_{M \rightarrow A} = \frac{R_A \sin 2\alpha}{(\sigma_2^A + R_A + R_A \cos 2\alpha) - (\sigma_2^A + R_A)} = \tan 2\alpha$$

Recall that we know the following values:

$$\sigma_1^A = 1100 \text{ MPa}$$

$$\sigma_2^A = 20 \text{ MPa}$$

$$\sigma_1^B = 1700 \text{ MPa}$$

$$\sigma_2^B = 40 \text{ MPa}$$

$$R_A = \frac{\sigma_1^A - \sigma_2^A}{2} = 540 \text{ MPa}$$

$$R_B = \frac{\sigma_1^B - \sigma_2^B}{2} = 830 \text{ MPa}$$

We can now solve $(*)$ and $(**)$ for μ and α :

$$\boxed{\mu \approx 2.65}$$

$$\boxed{\alpha \approx 1.39 \text{ radians} \approx 79.65 \text{ degrees}}$$

We can use some geometry to solve for C_0 (see the labels on the diagram):

$$\cos(2\alpha - \pi/2) = \frac{a}{h}$$

We know a so we can further say:

$$h = \frac{\sigma_2^A + R_A + R_A \cos 2\alpha}{\cos(2\alpha - \pi/2)} \approx 155 \text{ MPa}$$

Now:

$$b = h \cos 2\alpha \approx 145 \text{ MPa}$$

So:

$$\boxed{C_0 \approx 45.64 \text{ MPa}}$$

Now we suspect that a 3rd sample had a pre-existing crack. Then $C_0 = 0$ and we can calculate the two points where the failure criterion curve intersects the Mohr circle. For the 1st test: $72.6^\circ < \alpha < 86.7^\circ$. For the 2nd test: $74.0^\circ < \alpha < 85.3^\circ$. **N.B.** You can solve the whole problem by plotting Test 1 and Test 2 on graph paper!

Problem 6 (Strain accumulation at the San Andreas Fault)

We are given the deformation gradient tensor:

$$D = \begin{pmatrix} 0.15 & 0.24 \\ 0.00 & -0.15 \end{pmatrix}$$

We can decompose the deformation gradient tensor into a symmetric tensor and an antisymmetric tensor:

$$D = \begin{pmatrix} 0.15 & 0.12 \\ 0.12 & -0.15 \end{pmatrix} + \begin{pmatrix} 0.00 & 0.12 \\ -0.12 & 0.00 \end{pmatrix}$$

We know that the San Andreas Fault trends N65°W so let us rotate the strain tensor counterclockwise into that coordinate system:

$$\begin{aligned} \epsilon^* &= \begin{pmatrix} \cos 65^\circ & \sin 65^\circ \\ -\sin 65^\circ & \cos 65^\circ \end{pmatrix} \begin{pmatrix} 0.15 & 0.12 \\ 0.12 & -0.15 \end{pmatrix} \begin{pmatrix} \cos 65^\circ & \sin 65^\circ \\ -\sin 65^\circ & \cos 65^\circ \end{pmatrix}^{-1} \\ &= \begin{pmatrix} -0.004 & -0.192 \\ -0.192 & 0.004 \end{pmatrix} \end{aligned}$$

We see that the fault-shear strains (off-diagonal components) are rather large compared to the fault-normal strains (diagonal components). This is exactly what we expect for a strike-slip fault such as the San Andreas! The dilatation Δ equals the trace (sum of the diagonal elements) of ϵ . Thus we can see $\Delta = 0$.

Problem 7 (T+S Problem 3-19)

Equation (3-132) tells us the functional form $w(x)$ of the plate deflection:

$$w(x) = w_0 e^{-x/\alpha} \left(\cos \frac{x}{\alpha} + \sin \frac{x}{\alpha} \right)$$

To get to the bending moment M we need to calculate the second derivative of $w(x)$:

$$\frac{d^2 w}{dx^2} = \frac{2w_0 e^{-x/\alpha}}{\alpha^2} \left(\sin \frac{x}{\alpha} - \cos \frac{x}{\alpha} \right)$$

The bending moment M is proportional to the second derivative of $w(x)$:

$$M = -D \frac{d^2 w}{dx^2}$$

We want to calculate the maximum value of M so we must take a derivative and set it equal to zero:

$$\frac{dM}{dx} = \frac{4D}{\alpha^3} e^{-x/\alpha} \cos \frac{x}{\alpha} = 0$$

The above is true for:

$$x_{\max} = \pm \frac{\pi\alpha}{2}$$

Now we can plug back in for M_m and get equation (3-138):

$$M_{\max} \approx -0.416 \frac{Dw_0}{\alpha^2}$$

We can use the values given to estimate the maximum bending moment in the lithosphere:

$$M_{\max} \approx -1.6 \times 10^{17} \text{ N}$$

The maximum bending (fiber) stress σ_{xx}^{\max} is then given by equation (3-86):

$$\sigma_{xx}^{\max} = \pm \frac{6M}{h^2} \approx 8.1 \times 10^8 \text{ N/m}^2$$

The + corresponds to the tensile stress at the top of the plate and the – corresponds to the compressive stress at the bottom of the plate.

Problem 8 (T+S Problem 3-22)

We are told that the Amazon River Basin has a width $w = 400$ km. We are to model the basin as an elastic plate subject to a central line load (see Figures 3-29 and 3-30). We can use equation (3-135) to solve for the flexural parameter α :

$$\begin{aligned}x_b &= \pi\alpha \\ \frac{400 \text{ km}}{2} &= \pi\alpha \\ \alpha &\approx 64 \text{ km}\end{aligned}$$

We know that $(\rho_m - \rho_s) = 700 \text{ kg/m}^3$ so we can solve for the flexural rigidity D :

$$\begin{aligned}\alpha &= \left[\frac{4D}{(\rho_m - \rho_s)g} \right]^{1/4} \\ D &\approx 2.9 \times 10^{22} \text{ Nm}\end{aligned}$$

Now we can solve for the thickness T_e of the elastic lithosphere:

$$D = \frac{ET_e^3}{12(1 - \nu^2)}$$

$T_e \approx 17 \text{ km}$

Note the rather small value of T_e here.

Problem 9 (Dabbahu laccolith)

Start from the flexure equation:

$$D \frac{d^4 w}{dx^4} = q(x) - P \frac{d^2 w}{dx^2}$$

Note that $P = 0$ here. We can separate variables and integrate the equation four times to get:

$$w(x) = \frac{qx^4}{24D} + \alpha x^3 + \beta x^2 + \gamma x + \delta$$

The Greek letters are constants to be determined by four boundary conditions:

$$\begin{aligned}w(\pm L/2) &= 0 \\ \left. \frac{dw}{dx} \right|_{x=\pm \frac{L}{2}} &= 0\end{aligned}$$

The first two BCs tell us that $\alpha = \gamma = 0$. The second two BCs tell us:

$$\begin{aligned}\beta &= \frac{-qL^2}{48D} \\ \delta &= \frac{qL^4}{384D}\end{aligned}$$

So we get:

$$w(x) = \frac{qx^4}{24D} + \frac{qL^2 x^2}{48D} + \frac{qL^4}{384D}$$

If we define:

$$w_0 \equiv \frac{qL^4}{384D}$$

Then we can rewrite $w(x)$ as:

$$w(x) = w_0 \left(1 - \frac{8}{L^2} x^2 + \frac{16}{L^4} x^4 \right)$$

If we write $w(x)$ as $a - bx^2 + cx^4$ then we can do a polynomial fit for the Dabbahu data. The parameter a tells us a value for w_0 — I got $w_0 \approx 148 \text{ mm}$. We can use b and c to calculate L as follows:

$$\frac{b}{c} = \frac{L^2}{2}$$

$$\Rightarrow L \approx 45 \text{ km}$$

Now we can use equation (3-127) to calculate the flexural rigidity D :

$$D = \frac{gL^4(\rho_c - \rho_{\text{mag}})}{256}$$

$$\Rightarrow D \approx 6.5 \times 10^{19} \text{ Nm}$$

We can use our value of D to solve for the plate thickness h :

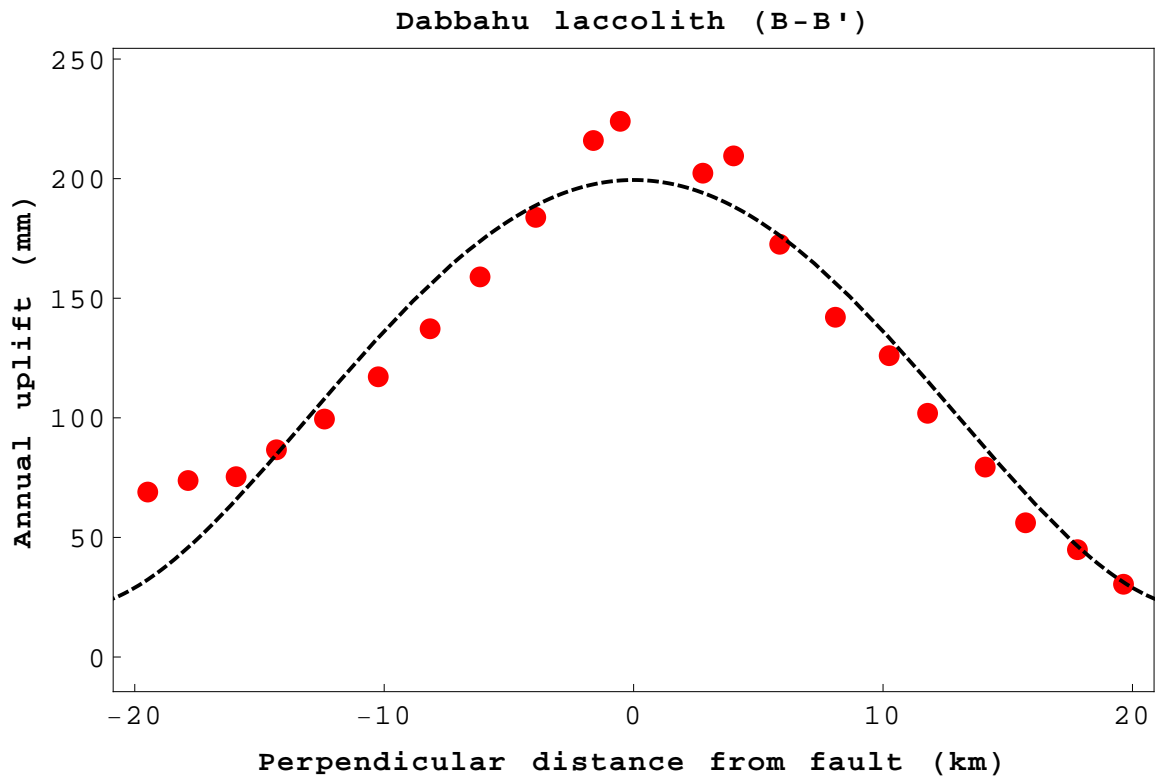
$$D = \frac{Eh^3}{12(1 - \nu^2)}$$

$$\Rightarrow h \approx 2.2 \text{ km}$$

Finally we can solve for the magma pressure p :

$$w_0 = \frac{q}{D} = \frac{\rho_c gh - p}{D}$$

$$\Rightarrow p \approx 57 \text{ MPa}$$



Two kilometers or so to the roof of the laccolith seems reasonable given that Dabbahu sits on a plate triple junction (Somalian-Nubian-Arabian).

We expect dykes to nucleate at areas of great tensional stress. We know how the normal strain $\epsilon_{xx}(x, y)$ relates to the normal stress $\sigma_{xx}(x, y)$:

$$\sigma_{xx}(x, y) = \frac{E}{1 - \nu^2} \epsilon_{xx}(x, y)$$

And we know how $\epsilon_{xx}(x, y)$ relates to the deflection $w(x)$:

$$\epsilon_{xx}(x, y) = -y \frac{d^2 w}{dx^2}$$

Now we can take the second derivative of the deflection $w(x)$ at evaluate that at $y = +h/2$ (the bottom of the plate):

$$\epsilon_{xx}(x) = \frac{qhL^2}{48D} \left(1 - \frac{12}{L^2} x^2 \right)$$

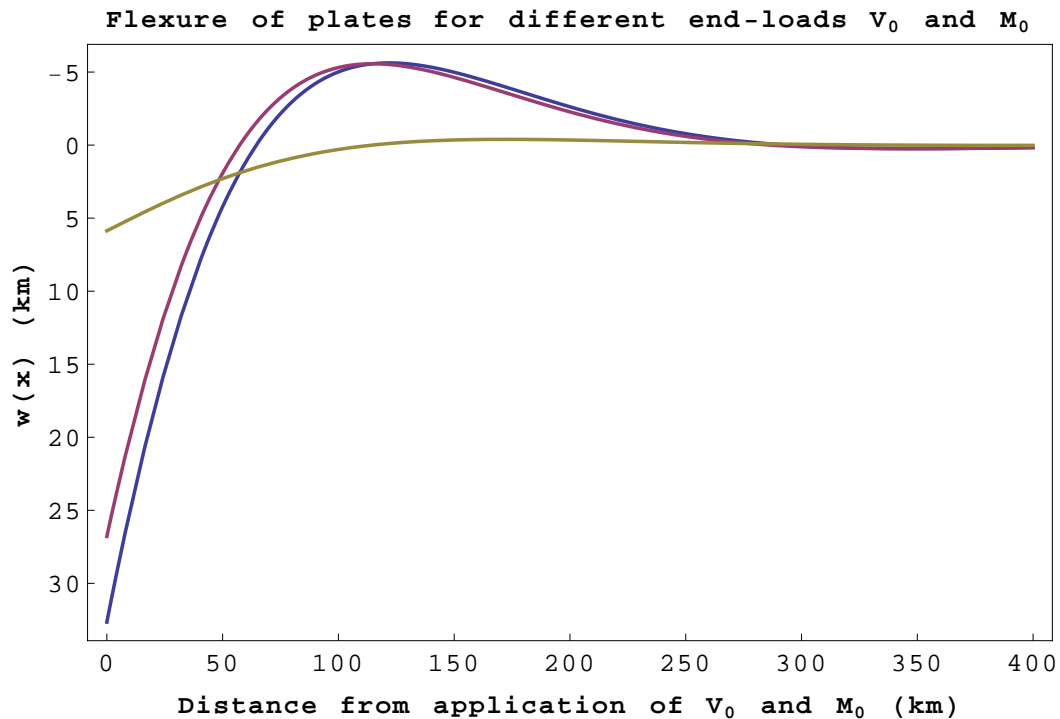
So the normal stresses $\sigma_{xx}(x)$ at the bottom of the plate can be found — recall that $D = Eh^3/12(1 - \nu^2)$:

$$\sigma_{xx}(x) = \frac{E}{1 - \nu^2} \epsilon_{xx}(x) = \frac{qL^2}{4h^2} \left(1 - \frac{12}{L^2} x^2 \right)$$

We want the value of $x \in [-L/2, L/2]$ that produces the greatest *negative* (tensional) value of stress. We can see that $\sigma_{xx}(x)$ plots as a concave-downward parabola. Thus the greatest negative stresses must occur at the endpoints of the model domain — at $x = -L/2$ and $x = +L/2$.

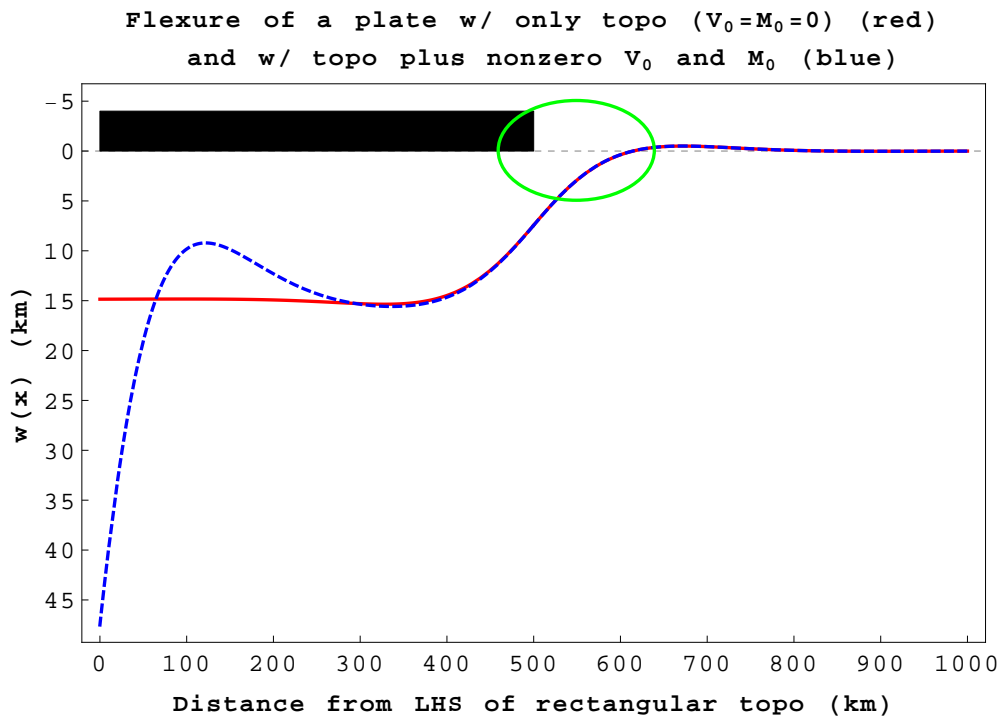
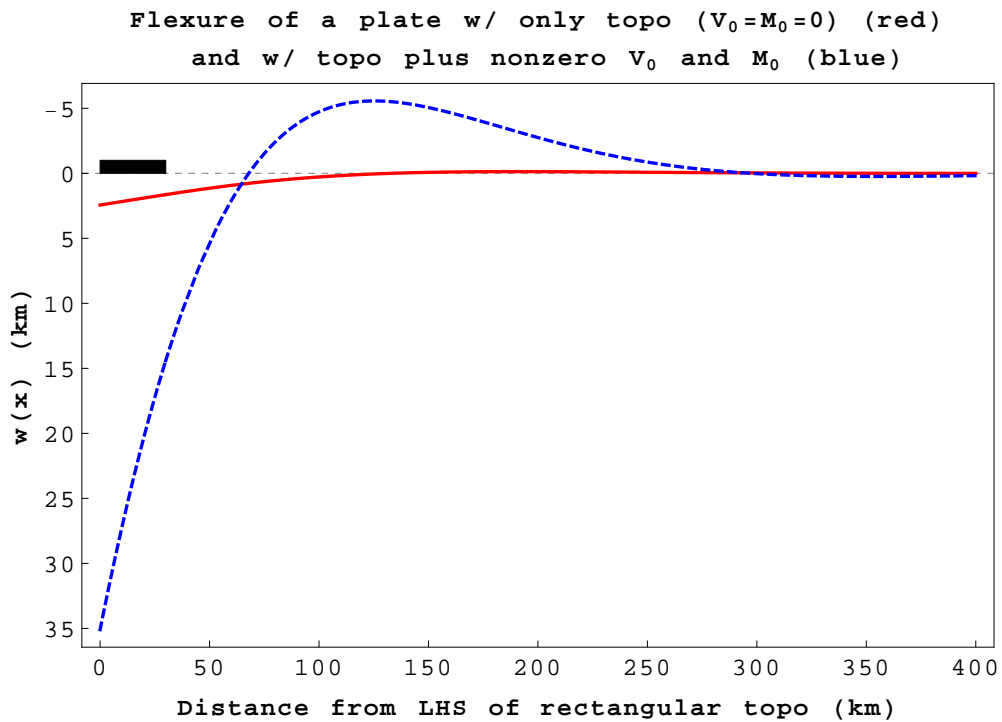
Problem 10 (Broken plate flexure)

Blue: $V_0 = -1.5 \times 10^{12}$ N/m and $M_0 = -5.0 \times 10^{17}$ N. Purple: $V_0 = 0$ N/m and $M_0 = -5.0 \times 10^{17}$ N. Yellow: $V_0 = -1.5 \times 10^{12}$ N/m and $M_0 = 0$ N.



The flexural response depends upon the flexural rigidity D of the plate which scales as the *cube*(!) of the elastic plate thickness. In other words: A modest change to T_e greatly changes the flexural response of the plate.

Problem 11 (Adding topography)



We *cannot* use the shape of the Ganges Foredeep Basin (circled in green) to study slab-pull on the subducting Indian Plate. This is because the flexural response of the plate around the Foredeep Basin looks the same with or without loading at $x = 0$. Thus topography alone controls the shape of the Basin — at least for the case at hand.